General Concepts

- Scalars, Vectors, Tensors, Matrices
- Fields
- Waves and wave equation
- Signal and Noise
- <u>Reading:</u>
 - > Telford et al., Sections A.2-3, A.5, A.7
 - Shearer, 2.1-2.2, 11.2, Appendix 2

Vector

Directional quantity

- Possesses 'amplitude' and 'direction' and nothing else...
 - Thus it can be described by its amplitude and two directional angles (*e.g.*, *azimuth* and *dip*).
- Characterized by projections on three selected axes: (x,y,z)...
 - ...plus an agreement that the projections are transformed appropriately whenever the frame of reference is rotated.



Tensor (informal)

- **Bi-Directional quantity**
 - 'Relationship' between two vectors;
 - Represented by a *matrix*:
 - 3×3 in three-dimensional space, 2×2 in two dimensions, etc.
 - ...this matrix, however, is transformed whenever the frame of reference is rotated.

Examples:

- Rotation operator, *R* in the plot below;
- Stress and strain in an elastic body.



Vector operations

Summation: c = a + b

 $c_x = a_x + b_x, c_y = a_y + b_y, c_z = a_z + b_z$ or simply: $c_i = a_i + b_i$

Scaling: $c = \lambda b$

$$c_x = \lambda b_x, c_y = \lambda b_y, c_z = \lambda b_z$$

$$c_i = \lambda b_i$$

Scalar (dot) product:

$$c = \mathbf{a} \cdot \mathbf{b} = a_x b_x + a_y b_y + a_z b_z$$

"Einstein's" notation: $c = a_i b_i$

Vector (cross) product:

$$c = \mathbf{a} \times \mathbf{b} = \begin{vmatrix} \mathbf{i} & \mathbf{j} & \mathbf{k} \\ a_x & a_y & a_z \\ b_x & b_y & b_z \end{vmatrix}$$
$$c_i = \epsilon_{iik} a_i b_k$$

Two key matrices

Unit (identity):

$$\boldsymbol{I} = \begin{vmatrix} 1 & 0 & 0 \\ 0 & 1 & 0 \\ 0 & 0 & 1 \end{vmatrix} \qquad \boldsymbol{I}_{ij} = \delta_{ij}$$

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 δ_{ii} is called the "Kronecker symbol":

$$\delta_{ij} = \begin{cases} 1 & \text{for } i = j \\ 0 & \text{for } i \neq j \end{cases}$$

Exercise: evaluate $\delta_{ii} = ?$

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Two key matrices

 Antisymmetric (or permutation, "Levi-Civita") symbol:

1 for (i, j, k)=even permutations of (1,2,3) ϵ_{ijk} =-1 for odd permutations of (1,2,3)0 otherwise

$$\epsilon_{ijk} =$$

key identities: $\epsilon_{ijj} \equiv \epsilon_{kik} \equiv \epsilon_{lli} \equiv 0$

vector cross-product: $c_i = \epsilon_{ijk} a_j b_k$

Exercise: evaluate $c_k = \epsilon_{ijk} \delta_{ij}$

Field

- Physical quantity which takes on values at a continuum of points in space and/or time
 - Represented by a function of coordinates and/or time:
 - > Scalar: f(x, y, z, t) or $f(\mathbf{r}, t)$
 - Examples: temperature, density, seismic velocity, pressure, gravity, electric potential
 - > Vector: $\boldsymbol{F}(\boldsymbol{r},t)$
 - Examples: particle displacement, velocity, or acceleration, force, electric or magnetic field, current

> Tensor

- 'relation' between two vectors
- Examples: strain and stress, electromagnetic field in electrodynamics
- The only way to describe *anisotropy*
- Always associated with some *source*, carries some kind of *energy*, and often able to propagate *waves*
- Everything in physics is fields!

Scalar Fields

Gradient

- Spatial derivative of a scalar field (say, temperature, T(x,y,z,t))
- It is a Vector field, denoted ∇T ('nabla' T):



Vector Fields Differential operations

Gradient of a vector field is a tensor: $(grad U_j)_i = \partial_i U_j$

Curl operation produces a new vector field:

$$(\operatorname{curl} \mathbf{F})_i = \epsilon_{ijk} \partial_j F_k$$

Two Important Relations

Divergence of a curl is always zero: $div(curl(\psi)) \equiv 0.$

This will be the S wave

• Curl of a gradient is zero: $curl(grad(\phi)) \equiv 0.$

This will be the P wave

These properties are easily verified using Einstein's notation (<u>try this!</u>):

 $(\operatorname{grad} U)_i = \partial_i U$ $(\operatorname{curl} \mathbf{F})_i = \epsilon_{iik} \partial_i F_k$

Static Fields and Waves

- Fields in geophysics typically exhibit either *static* or *wave* behaviours:
 - Static independent on time:

 $\frac{\partial T}{\partial t} = 0.$

 $\frac{1}{c^2}\frac{\partial^2 p}{\partial t^2}$

Stationary temperature distribution (geotherm).

 Wave – stable spatial pattern propagating with time:

 $\nabla^2 p = 0$ Acoustic (pressure) wave.

This is the typical form of wave equation; *c* is the velocity of propagation.

p = f(x - ct) Plane wave propagating along the X-axis.

> f(...) is the waveform at time t, it has its "zero" at x = ctand propagates to x > 0

Exercise: show this.

Signal and Noise

Geophysical data always contain some SIGNAL and some NOISE

- Signal 'deterministic' part that we want to know
 - Consistent with the method employed
- Noise anything else mixed into the measurement

Sources of noise:

- ♦Instrument
- Geologic sources
- Too simple theory (e.g., 2-D sounding in a 3-D Earth)
- Types of noise
 - Coherent (caused by the signal itself, worst of all)
 - Incoherent (random, coming from unrelated sources)
 - Such noise can be reduced by filtering

 Main task of data processing is to increase the signal/noise (S/N) ratio

S/N improvement by stacking

• "Stacking" (summation) is the most common approach to increasing the Signal/Noise ratio

 To derive the S/N improvement factor, consider stacking of N records with identical signals and random noise:

$$u_i(t) = s(t) + n_i(t)$$

• Stacked signal amplitude is proportional to N:

$$\sum_{i=1}^{N} u_i(t) = Ns(t) + \sum_{i=1}^{N} n_i(t)$$

Noise power increases ∞N (despite what it is commonly said, noise is not "attenuated" by stacking!):

$$\langle \left(\sum_{i=1}^{N} n_{i}(t)\right)^{2} \rangle = \sum_{i=1}^{N} \sum_{j=1}^{N} \langle n_{i}(t) n_{j}(t) \rangle = N \langle n^{2}(t) \rangle$$

• Therefore: $\frac{S}{N} = \sqrt{N} \frac{s}{n}$

S/N ratio increases as $\sqrt{\Lambda}$

Noise in Geophysical Measurements



- For seismics, the *signal* is represented by reflections and refractions
 - For 2D, also only those coming in-plane.
- Several factors cause degradation of the seismic signal:

