

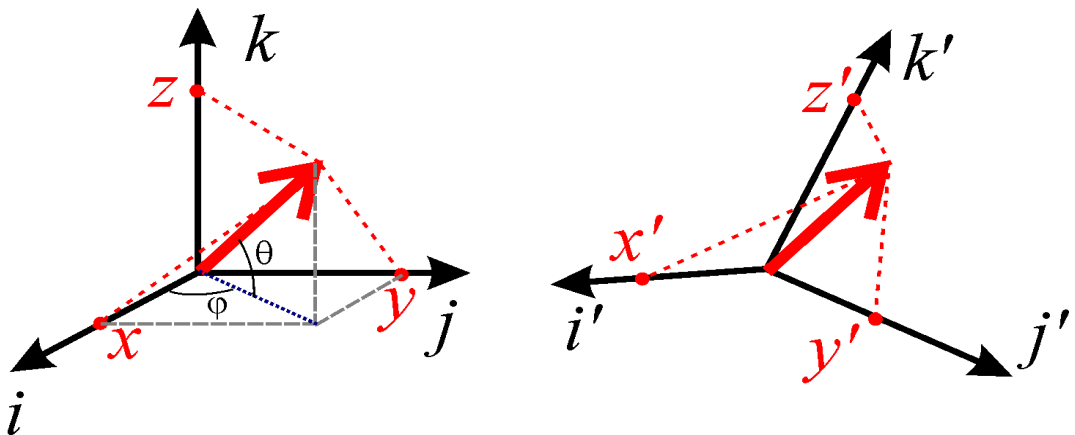
# General Concepts

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- Scalars, Vectors, Tensors, Matrices
- Fields
- Waves and wave equation
- Signal and Noise
- Reading:
  - › Telford *et al.*, Sections A.2-3, A.5, A.7
  - › Shearer, 2.1-2.2, 11.2, Appendix 2

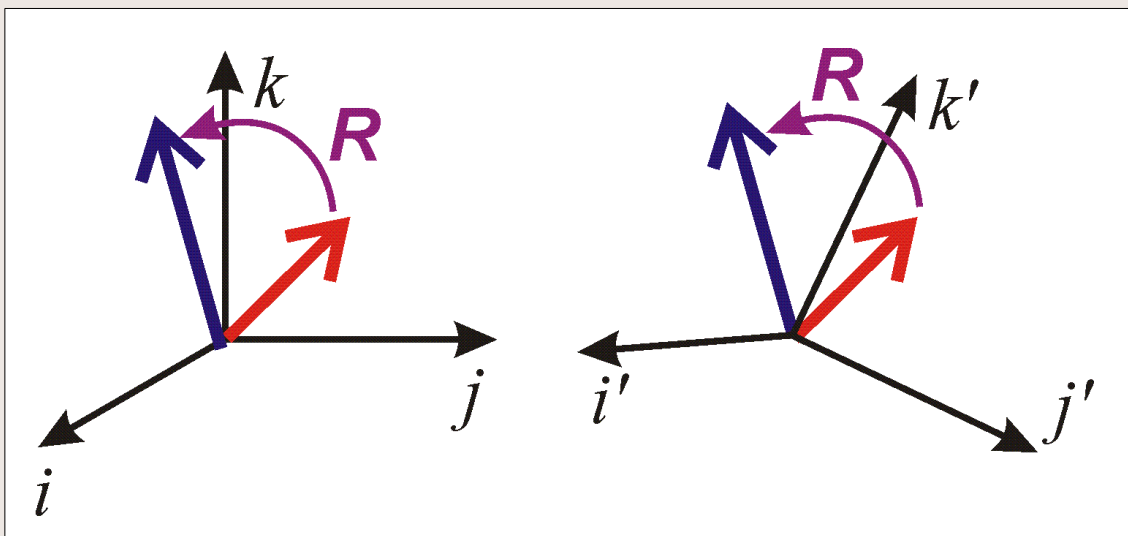
# Vector

- Directional quantity
  - ◆ Possesses 'amplitude' and 'direction' and nothing else...
    - Thus it can be described by its amplitude and two directional angles (*e.g.*, *azimuth* and *dip*).
  - ◆ Characterized by projections on three selected axes:  $(x,y,z)$ ...
    - ...plus an agreement that the projections are transformed appropriately whenever the **frame of reference** is rotated.



# Tensor (informal)

- Bi-Directional quantity
  - ♦ 'Relationship' between two vectors;
  - ♦ Represented by a *matrix*:
    - ♦  $3 \times 3$  in three-dimensional space,  $2 \times 2$  in two dimensions, etc.
    - ♦ ...this matrix, however, is transformed whenever the **frame of reference** is rotated.
- Examples:
  - ♦ Rotation operator,  $R$  in the plot below;
  - ♦ Stress and strain in an elastic body.



# Vector operations

- Summation:  $\mathbf{c} = \mathbf{a} + \mathbf{b}$

$$c_x = a_x + b_x, c_y = a_y + b_y, c_z = a_z + b_z$$

$$\text{or simply: } c_i = a_i + b_i$$

- Scaling:  $\mathbf{c} = \lambda \mathbf{b}$

$$c_x = \lambda b_x, c_y = \lambda b_y, c_z = \lambda b_z$$

$$c_i = \lambda b_i$$

- Scalar (dot) product:

$$c = \mathbf{a} \cdot \mathbf{b} = a_x b_x + a_y b_y + a_z b_z$$

$$\text{“Einstein’s” notation: } c = a_i b_i$$

- Vector (cross) product:

$$\mathbf{c} = \mathbf{a} \times \mathbf{b} = \begin{vmatrix} \mathbf{i} & \mathbf{j} & \mathbf{k} \\ a_x & a_y & a_z \\ b_x & b_y & b_z \end{vmatrix}.$$

$$c_i = \epsilon_{ijk} a_j b_k$$

# Two key matrices

- Unit (identity):

$$\mathbf{I} = \begin{bmatrix} 1 & 0 & 0 \\ 0 & 1 & 0 \\ 0 & 0 & 1 \end{bmatrix}$$

$$I_{ij} = \delta_{ij}$$

- ♦  $\delta_{ij}$  is called the “Kronecker symbol”:

$$\delta_{ij} = \begin{cases} 1 & \text{for } i = j \\ 0 & \text{for } i \neq j \end{cases}$$

**Exercise:** evaluate  $\delta_{ii} = ?$

# Two key matrices

- Antisymmetric (or permutation, “Levi-Civita”) symbol:

$$\epsilon_{ijk} = \begin{cases} 1 & \text{for } (i, j, k) = \text{even permutations of } (1, 2, 3) \\ -1 & \text{for odd permutations of } (1, 2, 3) \\ 0 & \text{otherwise} \end{cases}$$

$$\epsilon_{ijk} = \begin{matrix} & \begin{matrix} 0 & 1 & 0 \\ 0 & 0 & 1 \\ 0 & 1 & 0 \end{matrix} \\ \begin{matrix} 0 \\ 0 \\ 0 \end{matrix} & \begin{matrix} 0 & 1 & 0 \\ 0 & 0 & 1 \\ 0 & 1 & 0 \end{matrix} \\ \begin{matrix} 0 \\ -1 \\ 0 \end{matrix} & \begin{matrix} 0 & 1 & 0 \\ 0 & 0 & 1 \\ 0 & 1 & 0 \end{matrix} \end{matrix}$$

key identities:  $\epsilon_{ijj} \equiv \epsilon_{kik} \equiv \epsilon_{lli} \equiv 0$

vector cross-product:  $c_i = \epsilon_{ijk} a_j b_k$

Exercise: evaluate  $c_k = \epsilon_{ijk} \delta_{ij}$



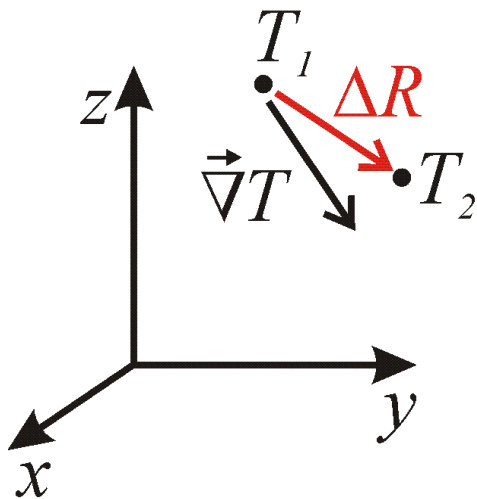
# Field

- Physical quantity which takes on values at a continuum of points in space and/or time
  - ◆ Represented by a function of coordinates and/or time:
    - Scalar:  $f(x, y, z, t)$  or  $f(\mathbf{r}, t)$ 
      - ◆ **Examples**: temperature, density, seismic velocity, pressure, gravity, electric potential
    - Vector:  $\mathbf{F}(\mathbf{r}, t)$ 
      - ◆ **Examples**: particle displacement, velocity, or acceleration, force, electric or magnetic field, current
    - Tensor
      - ◆ 'relation' between two vectors
      - ◆ **Examples**: strain and stress, electromagnetic field in electrodynamics
      - ◆ The only way to describe *anisotropy*
  - ◆ Always associated with some *source*, carries some kind of *energy*, and often able to propagate *waves*
- *Everything in physics is fields!*

# Scalar Fields

- Gradient

- ♦ Spatial derivative of a scalar field (say, temperature,  $T(x,y,z,t)$ )
- ♦ It is a Vector field, denoted  $\nabla T$  ('nabla' T):



$$\Delta T = T_2 - T_1$$

$$= \vec{\nabla} T \cdot \Delta \vec{R}$$

$$= \frac{\partial T}{\partial x} \Delta x + \frac{\partial T}{\partial y} \Delta y + \frac{\partial T}{\partial z} \Delta z$$

$$\Delta \vec{R} = \vec{i} \Delta x + \vec{j} \Delta y + \vec{k} \Delta z$$

$$\vec{\nabla} T = \vec{i} \frac{\partial T}{\partial x} + \vec{j} \frac{\partial T}{\partial y} + \vec{k} \frac{\partial T}{\partial z}$$



# Vector Fields

## *Differential operations*

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- **Gradient** of a vector field is a tensor:

$$(\text{grad } U_j)_i = \partial_i U_j$$

- **Curl** operation produces a new vector field:

$$(\text{curl } \mathbf{F})_i = \epsilon_{ijk} \partial_j F_k$$

# Two Important Relations

- Divergence of a curl is always zero:

$$\text{div}(\text{curl}(\boldsymbol{\psi})) \equiv 0.$$

This will be the S wave

- Curl of a gradient is zero:

$$\text{curl}(\text{grad}(\phi)) \equiv 0.$$

This will be the P wave

- These properties are easily verified using Einstein's notation (try this!):

$$(\text{grad } U)_i = \partial_i U$$

$$(\text{curl } \mathbf{F})_i = \epsilon_{ijk} \partial_j F_k$$

# Static Fields and Waves

- Fields in geophysics typically exhibit either *static* or *wave* behaviours:

- ♦ Static – independent on time:

$$\frac{\partial T}{\partial t} = 0. \quad \text{Stationary temperature distribution (geotherm).}$$

- ♦ Wave – stable spatial pattern propagating with time:

$$\frac{1}{c^2} \frac{\partial^2 p}{\partial t^2} - \nabla^2 p = 0 \quad \text{Acoustic (pressure) wave.}$$

This is the typical form of wave equation;  $c$  is the velocity of propagation.

$$p = f(x - ct) \quad \text{Plane wave propagating along the } X\text{-axis.}$$

$f(\dots)$  is the waveform at time  $t$ , it has its “zero” at  $x = ct$  and propagates to  $x > 0$

**Exercise:**  
show this.

# Signal and Noise

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- Geophysical data always contain some **SIGNAL** and some **NOISE**
  - ◆ Signal - 'deterministic' part that we want to know
    - Consistent with the method employed
  - ◆ Noise - anything else mixed into the measurement
- Sources of noise:
  - ◆ Instrument
  - ◆ Geologic sources
  - ◆ Too simple theory (e.g., 2-D sounding in a 3-D Earth)
- Types of noise
  - ◆ **Coherent** (caused by the signal itself, worst of all)
  - ◆ **Incoherent** (random, coming from unrelated sources)
    - Such noise can be reduced by filtering
- **Main task of data processing is to increase the signal/noise (S/N) ratio**

# $S/N$ improvement by stacking

- “Stacking” (summation) is the most common approach to increasing the Signal/Noise ratio
- To derive the  $S/N$  improvement factor, consider stacking of  $N$  records with identical signals and random noise:

$$u_i(t) = s(t) + n_i(t)$$

- Stacked **signal amplitude** is proportional to  $N$ :

$$\sum_{i=1}^N u_i(t) = Ns(t) + \sum_{i=1}^N n_i(t)$$

- **Noise power** increases  $\propto N$  (despite what is commonly said, noise is not “attenuated” by stacking!):

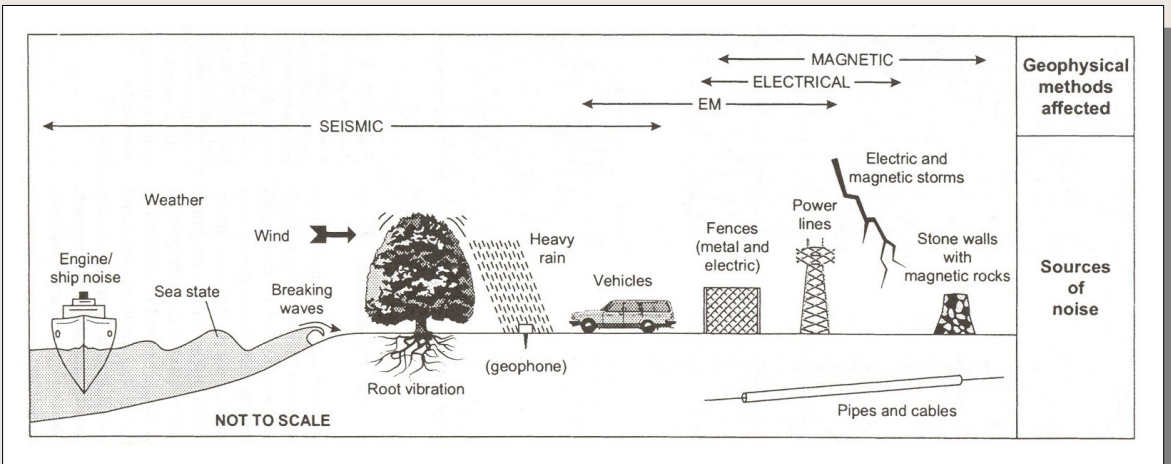
$$\langle \left( \sum_{i=1}^N n_i(t) \right)^2 \rangle = \sum_{i=1}^N \sum_{j=1}^N \langle n_i(t) n_j(t) \rangle = N \langle n^2(t) \rangle$$

- Therefore:  $\frac{S}{N} = \sqrt{N} \frac{s}{n}$

$S/N$  ratio  
increases as  $\sqrt{N}$



# Noise in Geophysical Measurements



- ◆ For seismics, the *signal* is represented by reflections and refractions
  - ◆ For 2D, also only those coming in-plane.
- ◆ Several factors cause degradation of the seismic signal:

