## Geol 335.3

## Lab \#6: Common mid-point method

In this lab, you will use simple Matlab or Octave simulations to study the principle of the common-midpoint (CMP) reflection method. Tools and functions from the previous labs will be useful in this exercise.

## Theory

Unlike the common-shot seismic data studied in the previous lab, CMP records are collected by moving both sources and receivers in opposite directions so that their midpoint remains constant while the source-receiver distance increases. CMP gathers are usually presented in the form of time-offset seismic sections.
The primary use of CMP gathers is for stacking velocity analysis. In its simplest form, velocity analysis is performed by trying multiple pairs of normal-incidence reflection times $\left(t_{0}\right)$ and stacking velocities $V$, and for each of them, calculating some "semblance" function representing the degree of signal coherence along the hyperbola. The pair $\left(t_{0}, V\right)$ giving the largest semblance identifies a reflector at depth $z=t_{0} V / 2$, with average velocity $V$ above it.
There are many ways to calculate the semblance function from seismic records. For example, the stack (summation) of signal powers along the reflection hyperbola can be used:

$$
\begin{equation*}
\text { Semblance }\left(t_{0}, V\right)=\frac{1}{N} \sum_{i=1}^{N} u_{i}^{2}\left(\sqrt{t_{0}^{2}+\left(\frac{x_{i}}{V}\right)^{2}}\right) \tag{1}
\end{equation*}
$$

where $u_{i}(t)$ is the signal it $i^{\text {th }}$ channel, and $x_{i}$ is the source-receiver offset at its location. In this lab, we use a different measure of semblance which provides smoother peaks:

$$
\begin{equation*}
\text { Semblance }\left(t_{0}, V\right)=\text { const } \times \operatorname{smooth}\left\{\left[\sum_{i=1}^{N} u_{i}\left(\sqrt{t_{0}^{2}+\left(\frac{x_{i}}{V}\right)^{2}}\right)\right]^{2}\right\} \text {. } \tag{2}
\end{equation*}
$$

This is a smoothed squared stack of the waveforms evaluated along the reflection hyperbola.

For a horizontal reflector, the stacking velocity equals the averaged (in the sense discussed in class) velocity above it. When reflector dip $\alpha$ is present, the stacking velocity increases:

$$
\begin{equation*}
V_{\text {with dip }}=\frac{V_{\text {no dip }}}{\cos \alpha} . \tag{3}
\end{equation*}
$$

## Code

In another copy of your model1.m file (from lab 5), rename sources to midpoints. For each midpoint number $n$, you can obtain a CMP section by using a small modification of function reflection (...) called reflection_CMP (...). In this function, source positions are variable for each receiver, so that the midpoint is fixed. For example, for the first midpoint (denoted midpoints (1) in your code), the resulting section will be produced by

```
sec = reflection_CMP(midpoints(1),receivers,layer, 2.0);
```

To plot this section, you will need to first calculate the relative source-receiver distances (verify that this formula is correct!):

```
offsets(1,:) = 2*(receivers - midpoints(1))
```

Then, the plot is obtained by:

$$
\text { plot_section(offsets } \left.(1,:), \text { sec,' Offset }(m)^{\prime}, ' b-'\right)
$$

To compute the velocity spectra in eq. (2), we provide function semblance (). Look into its code. Note that it is constructed very similarly to reflection(), by first initializing a blank semblance and then adding to it contributions from all traces using interpl(). The output of semblance () also represents a trace section, which can be plotted by plot_section():
plot_section(V, sec,'V(km/s)', 'b-'),
where V is the array of trial stacking velocities.

## Assignments

Download and unpack archive file lab6.zip. Start with file model1.m and following its examples and commentaries, do the following:

1. [5\%] Put three points (or more if you like) into array midpoints. Complete the for loop in the code to calculate the source-receiver offsets by formula (5) above.
2. [5\%] Execute the modified script model1.m and pick three horizontal reflecting boundaries. To pick any boundary, click at two points and then press spacebar. Place the first boundary layer1 in the bottom half of the section, and layer2 and layer3 at a shallow depth.
After you have picked the boundaries, you can comment out the interactive commands picklayer and uncomment the load command. This will re-load your picked boundaries next time you run the script.
3. [20\%] For one midpoint near the middle of the line, generate reflection sections (by using reflection_CMP ()) with different velocities. Try several values of velocity to see how they affect the shapes of reflections.
Note that not all combinations of velocities are possible. The relative variations of velocities should be smaller than those of layer depths, so that the zero-offset reflection travel times (seen at the apexes of reflection hyperbolas) should decrease from layer1 to layer 3. You can start from values 2.5, 2.0, and 1.5 and adjust them to achieve good images.

After modeling $\sec 1, \sec 2$, and $\sec 3$, sum them together:

$$
\text { section }=\sec +\sec 2+\sec 3
$$

This summation should illustrate how reflections from different depths overlap in real seismic data.
4. Note in the report how the times of reflections at zero source-receiver offsets $\left(t_{0}\right)$ correspond to the depths of reflectors, and the moveouts (slopes) of reflection hyperbolas correspond to the lower or higher velocities. Are the hyperbolas shifted up-dip as they were for common-shot recording in Lab 5?

The resulting section will be saved in modell.mat file, and the following velocity analysis steps you can do in script model2.m. This script will not have to re-compute the synthetics and will run faster.
5. [10\%] In script model2.m, compete the call to semblance (), which uses the CMP gather section to compute and plot the semblance spectrum (eq. 2) for stacking velocities $\mathrm{V}=0.7: 0.05: 3.0 \mathrm{~m} / \mathrm{ms}$ (note that these units are the same as $\mathrm{km} / \mathrm{s}$ ). Try denser spacings of velocities and larger upper limit to obtain better images.

In the resulting figure, you should see three subplots: the input reflection section, the semblance plotted as a seismic section (also using plot_section (...) function), and a color image of the semblance using imagesc () function. A reversed color map called 'hot' is used. Note this easy and handy way of plotting images in Matlab.
6. [10\%] Determine whether the velocities and $t_{0}$ set for the three reflectors are correctly determined by the peaks in the semblance plot. From the plots, describe the velocity resolution (width of velocity peaks) varies with $t_{0}$ and velocity.
7. [20\%] Run modell.m again and pick two points to make a dipping reflecting boundary close to the depth of the middle reflector in the preceding test. Plot the common-midpoint gather. How does it differ from the horizontal-reflector case? Is it shifted up-dip or down-dip (or not shifted)? What happens if the dip is changed?
8. [10\%] Plot the semblance spectrum for the dipping interface case. How do the optimal stacking velocity change? Compare the result to the prediction from formula (3).
9. [10\%] Summarize the differences of the horizontal and dipping reflector cases.

Next, look into and execute script model3.m. It is similar to model1.m and contains a (somewhat simplified) simulation of the whole zero-offset seismic section (ZOS). The ZOS represents the result of placing a source at each midpoint position and recording reflections on a receiver located at the same midpoint. This type of section represents the (almost) final output of reflection seismic imaging and approximates the geological or structural layering of the subsurface.
10. [10\%] Note the similarity in positions of seismic reflections (blue wiggles) with the depths of reflectors converted to time using relation $t_{0}=2 z / V$ (red lines). However, try noticing and comment on the differences between the positions and dips of reflectors and the resulting reflections. To see these differences clearer, try steep dips and shallow reflectors, maybe even a near-vertical reflector reaching close to the surface. Are the true reflectors steeper and shallower than reflections in the ZOS , or vice versa?

## Hand in:

Zipped directories containing:

1. Matlab codes ("m-files") which you have modified;
2. Discussion in a Word file including Screen captures or jpeg/PDF figures.
