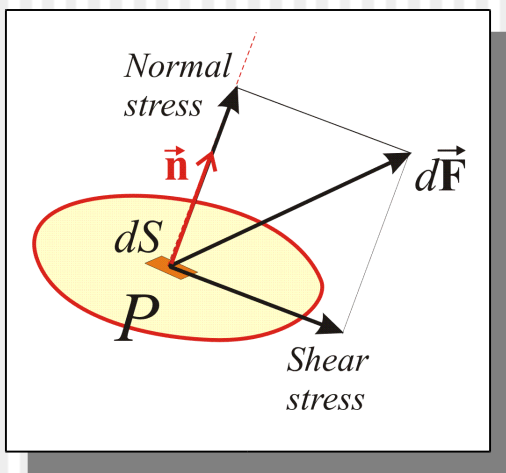


Theory of Elasticity

- Macroscopic theory
 - Equations of motion
 - Wave equations
 - Energy of a seismic wave
-
- Reading:
 - › Telford et al., Section 4.2
 - › Sheriff and Geldart, Sections 2.1-4

Stress

- Consider the interior of a deformed body:



At point P , force $d\vec{F}$ acts on any infinitesimal area dS . $d\vec{F}$ is a projection of *stress tensor*, σ , onto \mathbf{n} :

$$dF_i = \sigma_{ij} n_j dS$$

- Stress σ_{ij} is measured in [*Newton/m²*], or *Pascal* (unit of pressure).

- $d\vec{F}$ can be decomposed into two components relative to the orientation of the surface, \mathbf{n} :

- Parallel (*normal stress*)

$$(dF_n)_i = n_i \cdot (\text{projection of } \mathbf{F} \text{ onto } \mathbf{n}) = n_i \sigma_{kj} n_k n_j dS$$

- Tangential (*shear stress, traction*)

$$d\vec{F}_\tau = d\vec{F} - d\vec{F}_n$$

Note summation over k and j

Forces acting on a small cube

- Consider a small parallelepiped ($dx \times dy \times dz = dV$) within the elastic body.
- **Exercise 1:** show that the *force* applied to the parallelepiped from the outside is:

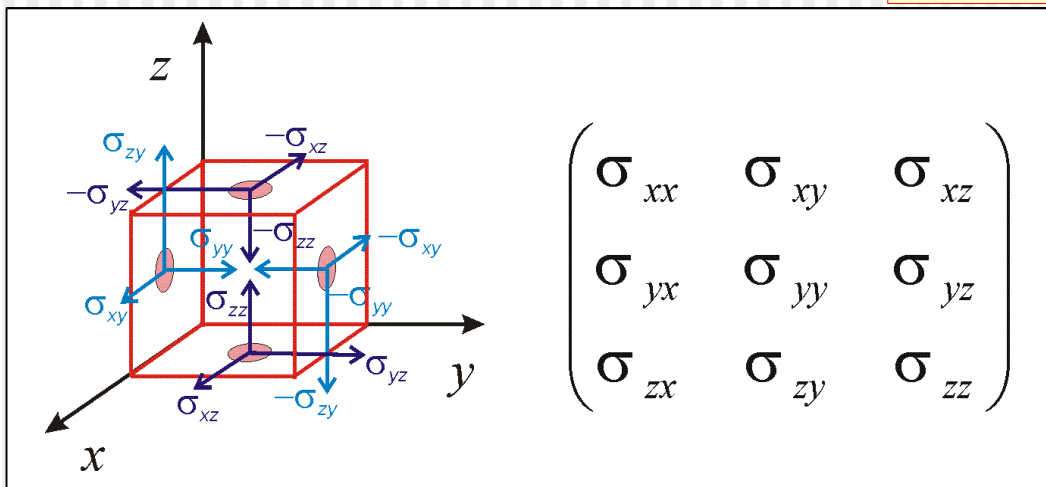
$$F_i = -\partial_j \sigma_{ij} dV$$

(This is simply minus divergence (convergence?) of stress!)

- **Exercise 2:** Show that *torque* applied to the cube from the outside is:

$$L_i = -\epsilon_{ijk} \sigma_{jk} dV$$

Again, keep in mind implied summations over repeated indices!



Symmetry of stress tensor

- Thus, L is proportional to dV : $L = O(dV)$
- The *moment of inertia* for any of the axes is proportional to $dV \cdot \text{length}^2$:

$$I_x = \int_{dV} (y^2 + z^2) \rho dV$$

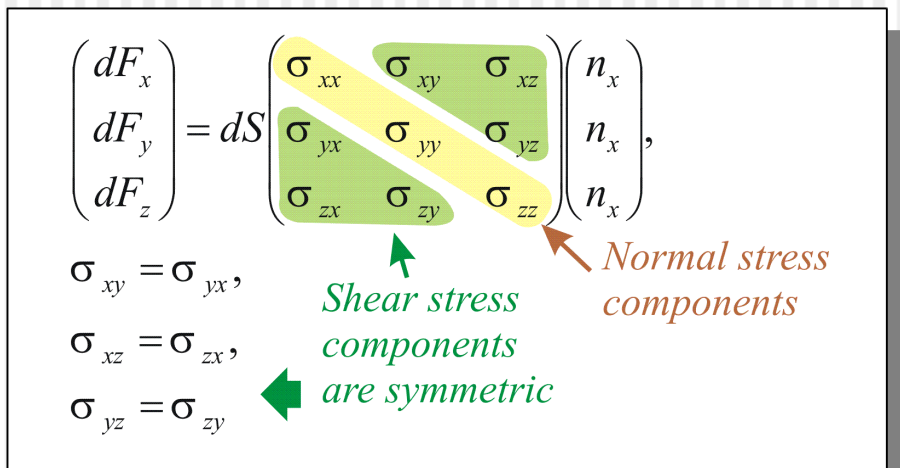
and so it tends to 0 faster than dV :

$$I = o(dV).$$

- Angular acceleration: $\theta = L/I$, must be finite as $dV \rightarrow 0$, and therefore:

$$L_i / dV = -\epsilon_{ijk} \sigma_{jk} = 0.$$

- Consequently, the stress tensor is *symmetric*: $\sigma_{ij} = \sigma_{ji}$
- σ_{ji} has only 6 independent parameters out of 9.



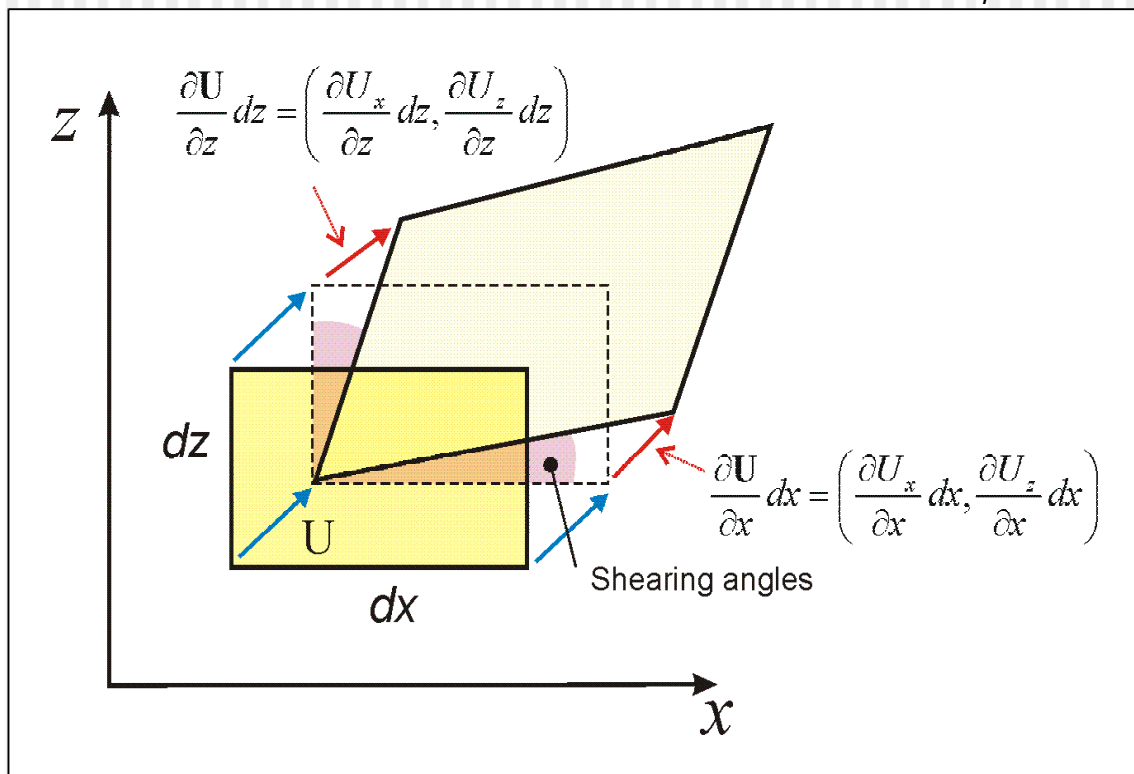
Strain

- Strain is a measure of deformation, i.e., *variation of relative displacement* as associated with a *particular direction* within the body
- It is, therefore, also a *tensor*
 - Represented by a matrix
 - Like stress, it is decomposed into *normal* and *shear* components
- Seismic waves yield strains of 10^{-10} - 10^{-6}
 - So we can rely on *infinitesimal* strain theory

Elementary Strain

- When a body is deformed, *displacements* (\mathbf{U}) of its points are dependent on (x, y, z) , and consist of:
 - Translation (**blue arrows** below)
 - Deformation (**red arrows**)

- Elementary strain is:
$$e_{ij} = \frac{\partial U_i}{\partial x_j}$$



Stretching and Rotation

- Exercise 1: Derive the elementary strain associated with unidirectional stretching of the body along the X axis:

$$\begin{pmatrix} x' \\ y' \end{pmatrix} = \begin{pmatrix} 1+\gamma & 0 \\ 0 & 1+\gamma \end{pmatrix} \begin{pmatrix} x \\ y \end{pmatrix}.$$

- Exercise 2: Derive the elementary strain associated with rotation by a small angle α :

$$\begin{pmatrix} x' \\ y' \end{pmatrix} = \begin{pmatrix} \cos \alpha & \sin \alpha \\ -\sin \alpha & \cos \alpha \end{pmatrix} \begin{pmatrix} x \\ y \end{pmatrix}.$$

Strain Components

- Anti-symmetric combinations of e_{ij} yield rotations of the body without changing its shape:
 - e.g., $\frac{1}{2}\left(\frac{\partial U_z}{\partial x} - \frac{\partial U_x}{\partial z}\right)$ yields rotation about the 'y' axis.
 - So, the case of $\frac{\partial U_z}{\partial x} = \frac{\partial U_x}{\partial z}$ is called *pure shear* (no rotation).
- To characterize *deformation*, only the symmetric component of the elementary strain is used:

$$\epsilon_{ij} = \frac{1}{2} \left(\frac{\partial U_i}{\partial x_j} + \frac{\partial U_j}{\partial x_i} \right),$$

$$\epsilon_{ij} = \epsilon_{ji}, \text{ where } i, j = x, y, \text{ or } z$$

$$\epsilon = \begin{pmatrix} \frac{\partial U_x}{\partial x} & \frac{1}{2} \left(\frac{\partial U_x}{\partial y} + \frac{\partial U_y}{\partial x} \right) & \frac{1}{2} \left(\frac{\partial U_x}{\partial z} + \frac{\partial U_z}{\partial x} \right) \\ \frac{1}{2} \left(\frac{\partial U_y}{\partial x} + \frac{\partial U_x}{\partial y} \right) & \frac{\partial U_y}{\partial y} & \frac{1}{2} \left(\frac{\partial U_y}{\partial z} + \frac{\partial U_z}{\partial y} \right) \\ \frac{1}{2} \left(\frac{\partial U_z}{\partial x} + \frac{\partial U_x}{\partial z} \right) & \frac{1}{2} \left(\frac{\partial U_z}{\partial y} + \frac{\partial U_y}{\partial z} \right) & \frac{\partial U_z}{\partial z} \end{pmatrix}$$

Dilatational Strain

(relative volume change during deformation)

- Original volume: $V = \delta x \delta y \delta z$
- Deformed volume: $V + \delta V = (1 + \varepsilon_{xx})(1 + \varepsilon_{yy})(1 + \varepsilon_{zz}) \delta x \delta y \delta z$
- Dilatational strain:

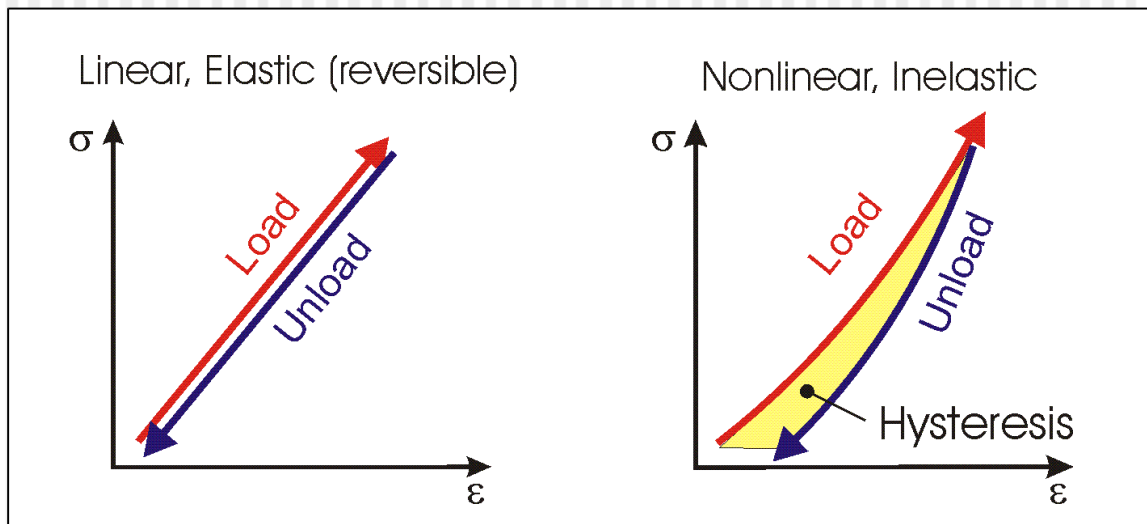
$$\Delta = \frac{\delta V}{V} = (1 + \varepsilon_{xx})(1 + \varepsilon_{yy})(1 + \varepsilon_{zz}) \approx \varepsilon_{xx} + \varepsilon_{yy} + \varepsilon_{zz}$$

$$\Delta = \varepsilon_{ii} = \partial_i U_i = \vec{\nabla} \cdot \vec{U} = \text{div } \vec{U}$$

- Note that *shearing strain does not change the volume.*

Hooke's Law (general)

- Hooke's law describes the **stress** developed in a **deformed body**:
 - $\mathbf{F} = -k\mathbf{x}$ for an ordinary spring (1-D)
 - $\sigma \sim \epsilon$ (in some sense) for a '*linear*', '*elastic*' 3-D solid. This is what these terms mean:



- For a general (*anisotropic*) medium, there are 36 coefficients of proportionality between six independent σ_{ij} and six ϵ_{ij} :

$$\sigma_{ij} = \Lambda_{ij,kl} \epsilon_{kl}.$$

Hooke's Law (isotropic medium)

- For isotropic medium, the strain/stress relation is described by just 2 constants:

$$\sigma_{ij} = \lambda \Delta \delta_{ij} + 2\mu \varepsilon_{ij}$$

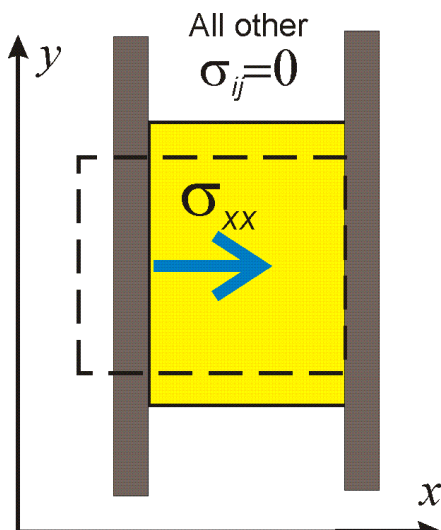
- δ_{ij} is the “Kronecker symbol” (unit tensor) equal 1 for $i=j$ and 0 otherwise;
 - λ and μ are called the *Lamé constants*.
- Although λ and μ provide a natural mathematical parametrization for $\sigma(\varepsilon)$, they are always intermixed in experimental environments.
 - Their combinations, called “*elastic modulae*” are typically directly measured.
 - Question: what are the units for λ and μ ?

Four Elastic Modulae

- Depending on boundary conditions (i.e., experimental setup), different combinations of λ' and μ may be convenient. These combinations are called *elastic constants*, or *modulae*:

- **Young's modulus and Poisson's ratio:**

- Consider a cylindrical sample uniformly squeezed along axis X:



All other $\sigma_{ij} = 0$

$$\sigma_{xx} = \lambda' \Delta + 2\mu \epsilon_{xx},$$

$$\sigma_{yy} = \lambda' \Delta + 2\mu \epsilon_{yy} = 0,$$

$$\sigma_{zz} = \lambda' \Delta + 2\mu \epsilon_{zz} = 0 \Rightarrow \epsilon_{yy} = \epsilon_{zz} = \frac{-\lambda' \Delta}{2\mu}.$$

Young's modulus: $E = \frac{\sigma_{xx}}{\epsilon_{xx}} = \frac{2\mu(3\lambda' + 2\mu)}{\lambda' + \mu}$

Poisson's ratio: $\nu = -\frac{\epsilon_{zz}}{\epsilon_{xx}} = \frac{\lambda'}{2(\lambda' + \mu)}$

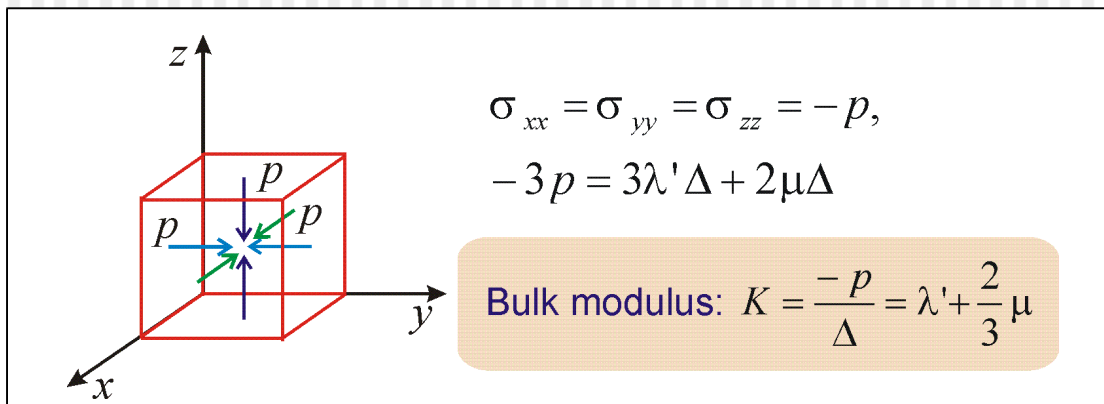
- Note: The Poisson's ratio is more often denoted σ
- It measures the ratio of λ and μ :

$$\frac{\mu}{\lambda} = \frac{1}{2\sigma} - 1$$

Four Elastic Modulae

◆ Bulk modulus, K

- Consider a cube subjected to hydrostatic pressure



- ◆ The constant μ complements K in describing the shear rigidity of the medium. Thus, μ is also called the '*rigidity modulus*'
- ◆ For rocks:
 - Generally, $10 \text{ Gpa} < \mu < K < E < 200 \text{ Gpa}$
 - $0 < \nu < \frac{1}{2}$ always; for rocks, $0.05 < \nu < 0.45$, for most, ν is near 0.25.
- ◆ For liquids, $\nu = \frac{1}{2}$ and $\mu = 0$ (no shear resistance)

Equations of Motion

(Motion of the elastic body with time)

- Uncompensated net force will result in *acceleration* (Newton's law):

Newton's law:
$$\rho \delta V \frac{\partial^2 U_i}{\partial t^2} = F_i$$

$$\rho \frac{\partial^2 U_i}{\partial t^2} = \left(\frac{\partial \sigma_{ix}}{\partial x} + \frac{\partial \sigma_{iy}}{\partial y} + \frac{\partial \sigma_{iz}}{\partial z} \right)$$

$$\begin{aligned} \rho \frac{\partial^2 U_x}{\partial t^2} &= \frac{\partial}{\partial x} \left(\lambda' \Delta + 2\mu \frac{\partial U_x}{\partial x} \right) + \mu \frac{\partial}{\partial y} \left(\frac{\partial U_x}{\partial y} + \frac{\partial U_y}{\partial x} \right) + \mu \frac{\partial}{\partial z} \left(\frac{\partial U_x}{\partial z} + \frac{\partial U_z}{\partial x} \right) \\ &= \lambda' \frac{\partial \Delta}{\partial x} + \mu \frac{\partial}{\partial x} \left(\frac{\partial U_x}{\partial x} + \frac{\partial U_y}{\partial y} + \frac{\partial U_z}{\partial z} \right) + \mu \left(\frac{\partial^2 U_x}{\partial x^2} + \frac{\partial^2 U_x}{\partial y^2} + \frac{\partial^2 U_x}{\partial z^2} \right) \\ &= (\lambda' + \mu) \frac{\partial \Delta}{\partial x} + \mu \nabla^2 U_x \end{aligned}$$

- Therefore, the *equations of motion* for the components of \mathbf{U} :

$$\rho \frac{\partial^2 U_i}{\partial t^2} = (\lambda + \mu) \frac{\partial \Delta}{\partial x_i} + \mu \nabla^2 U_i$$

Wave potentials

Compressional and Shear waves

- These equations describe two types of waves.
- The general solution has the form (“Lame theorem”):

$$\vec{U} = \vec{\nabla} \phi + \vec{\nabla} \times \vec{\psi}. \quad (\text{or } U_i = \partial_i \phi + \epsilon_{ijk} \partial_j \psi_k)$$

$$\vec{\nabla} \cdot \vec{\psi} = 0.$$

Because there are 4 components in ψ and ϕ only 3 in \mathbf{U} , we need to constrain ψ .

- Exercise: substitute the above into the equation of motion:

$$\rho \frac{\partial^2 U_i}{\partial t^2} = (\lambda + \mu) \frac{\partial \Delta}{\partial x_i} + \mu \nabla^2 U_i$$

and show:

$$\rho \frac{\partial^2 \phi}{\partial t^2} = (\lambda + 2\mu) \nabla^2 \phi, \quad \leftarrow \text{P-wave (scalar) potential.}$$

$$\rho \frac{\partial^2 \psi_i}{\partial t^2} = \mu \nabla^2 \psi_i, \quad \leftarrow \text{S-wave (vector) potential.}$$

Wave velocities

Compressional and Shear waves

- These are wave equations; compare to the general form of equation describing wave processes:

$$\left[\frac{1}{c^2} \frac{\partial^2}{\partial t^2} - \nabla^2 \right] f(x, y, z, t) = 0$$

- *Compressional* (*P*) wave velocity:

$$v_P = \sqrt{\frac{\lambda + 2\mu}{\rho}} = \sqrt{\frac{K + \frac{4}{3}\mu}{\rho}}$$

- *Shear* (*S*) wave velocity:

- $V_S < V_P$

- for $\sigma = 0.25$: $V_P/V_S = \sqrt{3}$

$$v_S = \sqrt{\frac{\mu}{\rho}}$$

- Note that the V_P/V_S depends on the Poisson ratio alone:

$$\frac{V_P}{V_S} = \sqrt{\frac{\mu}{\lambda + 2\mu}} = \sqrt{\frac{1/2 - \sigma}{1 - \sigma}}$$

Notes on the use of potentials

- Wave potentials are very useful for solving elastic wave problems
- Just take ϕ or ψ satisfying the wave equation, e.g.:

$$\phi(\vec{r}, t) = A e^{i\omega(t - \frac{\vec{r}\vec{n}}{V_P})} \quad \text{(plane wave)}$$

...and use the equations for potentials to derive the displacements:

$$\vec{U} = \vec{\nabla}\phi + \vec{\nabla}\times\vec{\psi}.$$

...and stress from Hooke's law:

$$\sigma_{ij} = \lambda\Delta\delta_{ij} + 2\mu\epsilon_{ij}$$

- *Displacement amplitude =*
 $\omega \times (\text{potential amplitude}) / V$
- *Power = $\frac{1}{2}\rho[\omega^2 \times (\text{potential amplitude}) / V]^2$*

This is *velocity* amplitude

Example:

Compressional (P) wave

- Scalar potential for *plane harmonic* wave:

$$\phi(\vec{r}, t) = A e^{i\omega(t - \frac{\vec{r}\vec{n}}{V_P})}$$

- Displacement:

$$u_i(\vec{r}, t) = \partial_i \phi(\vec{r}, t) = \frac{-i\omega n_i}{V_P} A e^{i\omega(t - \frac{\vec{r}\vec{n}}{V_P})}$$

note that the *displacement is always along \mathbf{n}* .

- Strain:

$$\varepsilon_{ij}(\vec{r}, t) = \partial_i u_j(\vec{r}, t) = \frac{\omega^2 n_i n_j}{V_P^2} A e^{i\omega(t - \frac{\vec{r}\vec{n}}{V_P})}$$

- Dilatational strain:

$$\Delta = \varepsilon_{ii}(\vec{r}, t) = \frac{\omega^2}{V_P^2} A e^{i\omega(t - \frac{\vec{r}\vec{n}}{V_P})} = \frac{\omega^2}{V_P^2} \phi(\vec{r}, t)$$

- Stress:

$$\sigma_{ij}(\vec{r}, t) = \frac{\omega^2}{V_P^2} (\lambda \delta_{ij} + 2\mu n_i n_j) \phi(\vec{r}, t)$$

- Question: what wavefield would we have if used $\cos()$ or $\sin()$ function instead of complex $\exp()$ in the expression for potential above?

Impedance

- *Impedance, Z* , is a measure of the amount of resistance to particle motion.
- In elasticity, impedance is a *ratio of stress to particle velocity*.
 - Thus, for a given applied stress, particle velocity is inversely proportional to impedance.
 - For P wave, in the direction of its propagation:

$$Z(\vec{r}, t) = \frac{\sigma_{nn}(\vec{r}, t)}{\dot{u}_n(\vec{r}, t)} = \frac{\lambda + 2\mu}{V_P} = \rho V_P.$$

- impedance does not depend on frequency but *depends on the wave type and incidence direction*.

Elastic Energy Density

- Mechanical work is required to deform an elastic body; as a result, elastic energy is accumulated in the strain/stress field
- When released, this energy gives rise to earthquakes and seismic waves
- For a loaded spring (1-D elastic body),
 $E = \frac{1}{2}kx^2 = \frac{1}{2}Fx$
- Similarly, for a deformed elastic medium, *energy density* is:

$$E = \frac{1}{2} \sigma_{ij} \epsilon_{ij}$$

Elastic Energy Density *in a plane wave*

- For a plane wave:

$$u_i = u_i(t - \vec{p} \cdot \vec{x})$$

$$\varepsilon_{ij} = \frac{1}{2}(\partial_i u_j + \partial_j u_i) = -\frac{1}{2}(\dot{u}_i p_j + \dot{u}_j p_i).$$

- ...and therefore:

$$\frac{1}{2} \sigma_{ij} \varepsilon_{ij} = \frac{1}{2} [(\lambda + \mu)(\vec{p} \cdot \vec{u})^2 + \mu(\vec{u} \cdot \vec{u})(\vec{p} \cdot \vec{p})]$$

- For P - and S -waves, this gives:

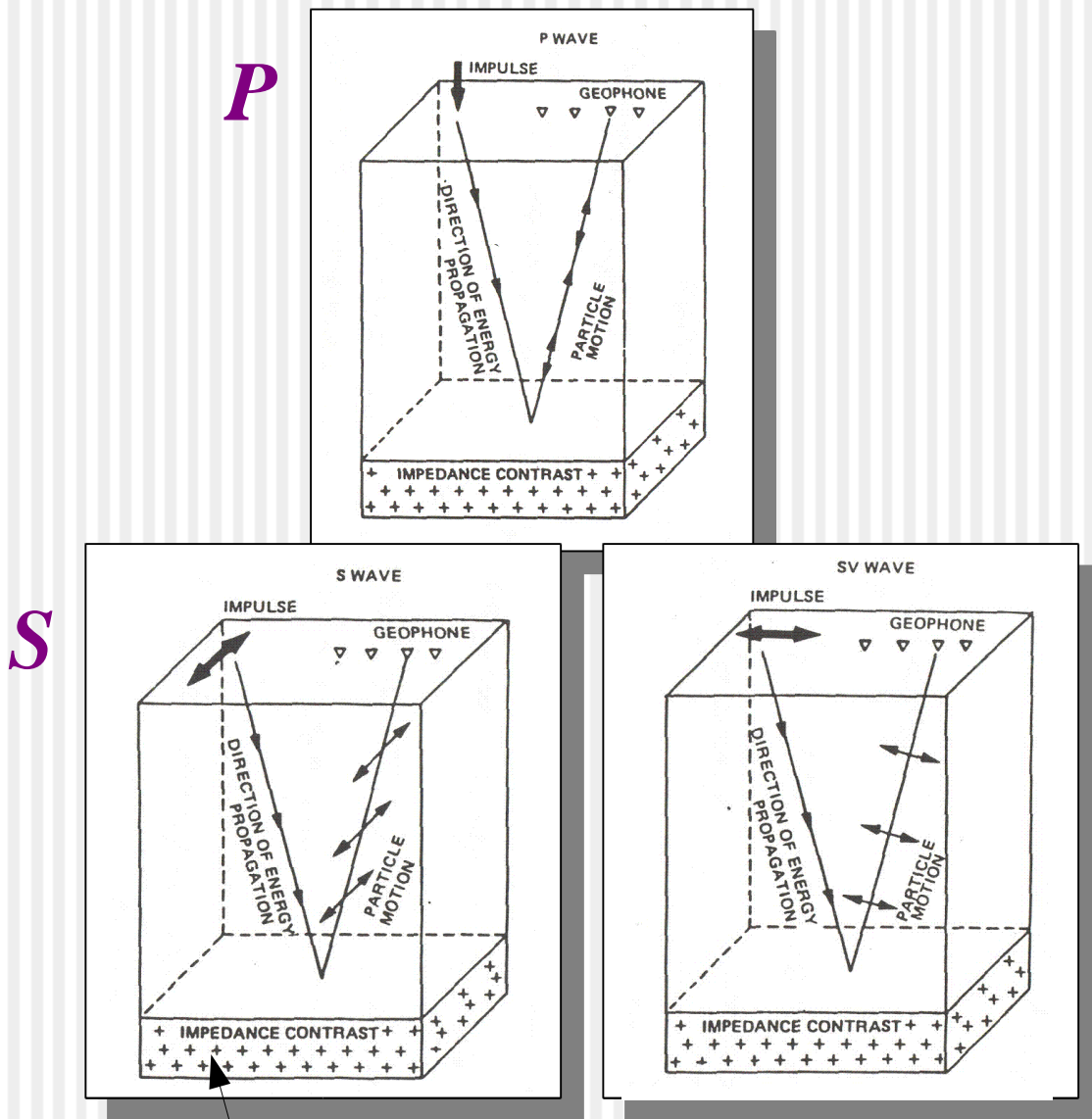
$$\frac{1}{2} \sigma_{ij} \varepsilon_{ij} = \frac{1}{2} (\lambda + 2\mu) p^2 \vec{u}^2 = \frac{1}{2} \rho \dot{\vec{u}}^2 \quad \text{P-wave}$$

$$\frac{1}{2} \sigma_{ij} \varepsilon_{ij} = \frac{1}{2} (\mu) p^2 \vec{u}^2 = \frac{1}{2} \rho \dot{\vec{u}}^2 \quad \text{S-wave}$$

- Thus, in a wave, *strain energy equals the kinetic energy* Energy is *NOT* conserved locally!
- *Energy propagates at the same speed* as the wave pulse

Wave Polarization

- Elastic solid supports two types of *body waves*:



Note that this is still an **ISOTROPIC** reflector. In general, reflection will intermix the S-wave polarization modes, and P-wave will convert into SV upon reflection.