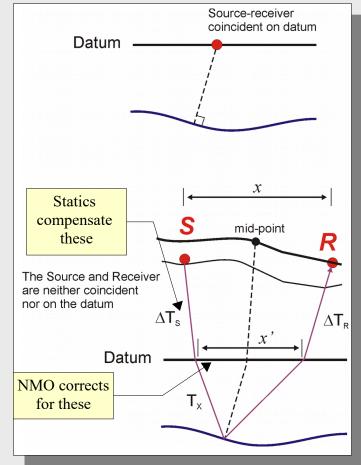
Geometrical Seismics *Reflection*

- Normal moveout (NMO)
- Normal moveout correction
- Dip moveout (DMO)
- <u>Reading:</u>
 - Sheriff and Geldart, Chapter 4.1.

Zero-Offset Section (The goal of reflection imaging)

- The Ideal of reflection imaging is sources and receivers collocated on a flat horizontal surface ("datum").
- In reality, however, we record at sourcereceiver offsets, and over complex topography.
- Two types of corrections are applied to compensate these factors:

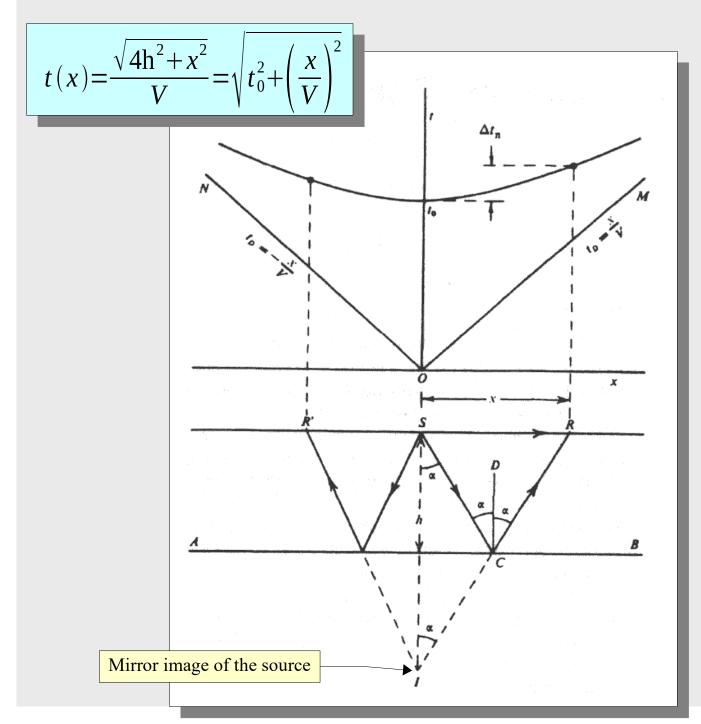


- Statics "place" sources and receivers onto the datum;
- Normal Moveout Corrections "transforms" the records into as if they were recorded at collocated sources and receivers.
- As a result of these corrections (plus stacking to attenuate noise), we obtain a zero-offset section.

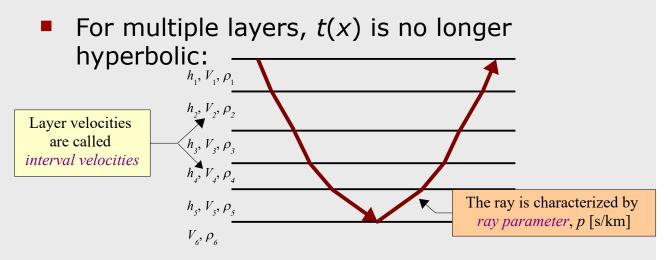
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Normal moveout New!

- Symmetrical hyperbola
- Reflected rays propagate as if from a source at depth



Reflection travel-times (*Multiple layers*)



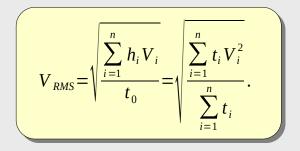
For practical applications (near-vertical incidence, $pV_i <<1$), t(x) still can be approximated as:

$$x_n(p) = \sum_{i=1}^n \frac{h_i p V_i}{\sqrt{1 - (pV_i)^2}} \approx p \sum_{i=1}^n h_i V_i [1 + \frac{1}{2} (pV_i)^2] \approx p \sum_{i=1}^n h_i V_i,$$

hence: $p = \frac{x_n(p)}{\sum_{i=1}^n h_i V_i}$, $t_n(p) = \sum_{i=1}^n \frac{h_i}{V_i \sqrt{1 - (pV_i)^2}} \approx \sum_{i=1}^n \frac{h_i}{V_i} [1 + \frac{1}{2} (pV_i)^2] = t_0 + \frac{1}{2} p^2 \sum_{i=1}^n h_i V_i$

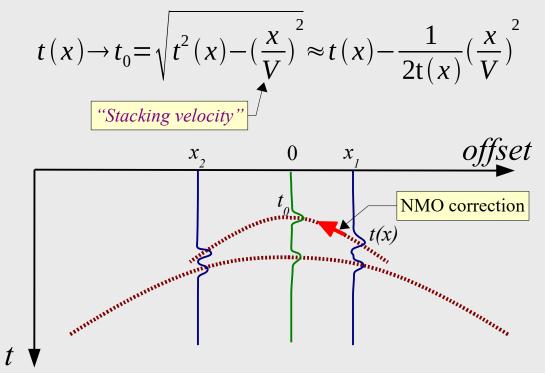
$$t_n(x) \approx t_0 + \frac{1}{2t_0} \left(\frac{x}{V_{RMS}}\right)^2$$

here, V_{RMS} is the RMS (root-mean-square) velocity:



Normal Moveout (NMO) correction

NMO correction transforms a reflection record at offset x into a normal-incidence (x=0) record:

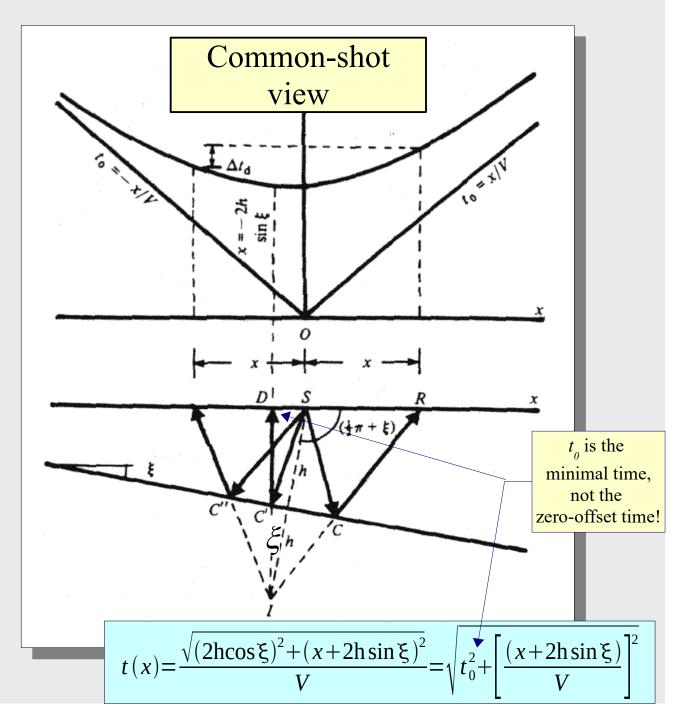


- Stacking velocity is determined from the data, as a parameter of the reflection hyperbola that is best aligned with the reflection event
- Note that NMO correction affects the shallower and slower reflections stronger
 - This is called "NMO stretching"

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Dipping reflector New!

- Hyperbola of the same shape but with the apex shifted up-dip
- Asymptotically the same moveouts





Dip moveout

For small offsets (x << h) and dips ($h \sin \xi << x$):

$$t(x) = \sqrt{t_0^2 + \left[\frac{(x+2h\sin\xi)}{V}\right]^2} \approx t_0 \left[1 + \frac{x^2 + 4hx\sin\xi}{2(t_0V)^2}\right].$$

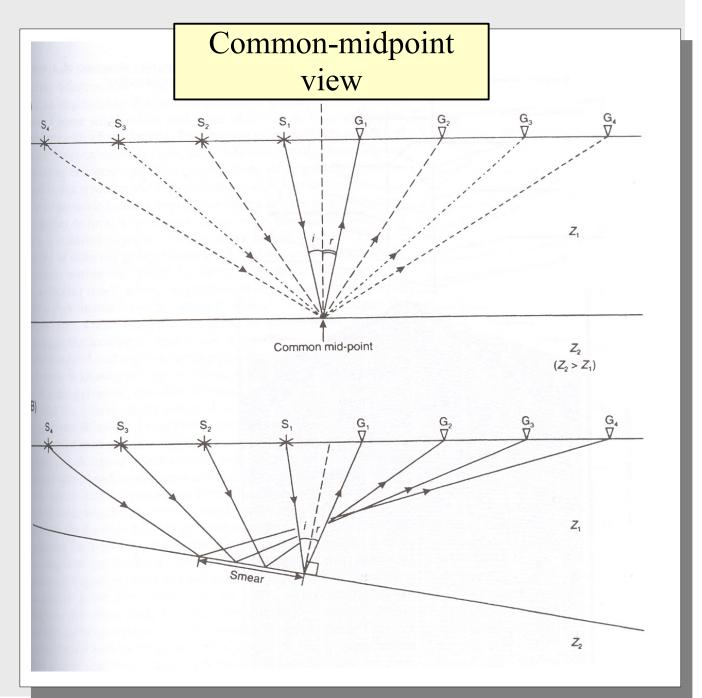
$$t(x) \approx t_0 + \frac{1}{2t_0 V^2} + \frac{1}{V}$$
Apex \approx Zero-offset time
Normal moveout term

 Reflector dip ξ can be measured from the *dip moveout*:
 This ratio is

$$\sin \xi \approx \frac{V}{2} \frac{t(x) - t(-x)}{x} \equiv \frac{V}{2} \frac{t_{Downdip} - t_{Updip}}{x}.$$

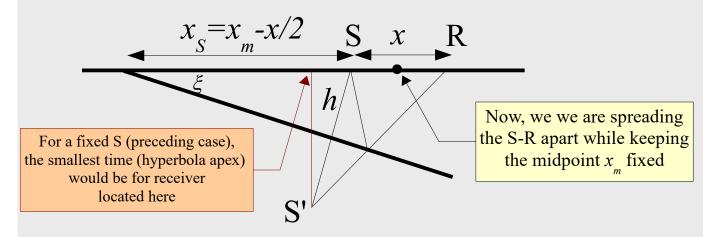
Dip moveout in CMP gathers

- The travel-time hyperbola becomes symmetrical
- Reflection points are *smeared* up-dip with increasing offset
- Asymptotic velocities are greater than the true velocity



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Stacking velocity in New! the presence of dip



 For a fixed x_m, the dependence of the S-R time on the offset x is

$$t(x) = \frac{1}{V} \sqrt{(x+2h\sin\xi)^2 + (2h\cos\xi)^2}$$

$$t(x) = \frac{1}{V} \sqrt{[x+(2x_m-x)\sin^2\xi]^2 + [(2x_m-x)\sin\xi\cos\xi]^2}$$

$$t(x) = \frac{1}{V} \sqrt{(2x_m\sin\xi)^2 + (x\cos\xi)^2}$$

continued...

CMP Stacking velocity in the presence of dip (*cont.*)

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This equation describes a hyperbola similar to the NMO equation (compare to: $t_{NMO}(x) = \sqrt{t_0^2 + \left(\frac{x}{V}\right)^2}$):

$$t(x) = \frac{1}{V} \sqrt{\left(2x_m \sin \xi\right)^2 + \left(x \cos \xi\right)^2} = \sqrt{\left(\frac{2x_m \sin \xi}{V}\right)^2 + \left(\frac{x \cos \xi}{V}\right)^2}$$
Zero-offset time
Hyperbolic moveout

Thus, because of the dip, the effective velocity is increased:

$$V_{Dip} = \frac{V}{\cos \xi}.$$

- This means that when stacking velocities are measured from a CMP gather, dipping reflectors will result in higher velocities (flatter reflection hyperbola)
- As a result, reflectors with conflicting dips cannot be NMO-corrected and stacked accurately.
 - Processing step called *DMO* corrects this problem.