### Geometrical Seismics *Basic concepts of Reflection imaging*

- Normal moveout (NMO)  $\bullet$
- Normal moveout correction  $\bullet$
- Dip moveout (DMO)  $\bullet$
- . Reading:
	- ➢ Sheriff and Geldart, Chapter 4.1

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# Zero-Offset Section

(Ideal picture and of reflection imaging and the goal of "CMP stacking")

- **The Ideal of reflection imaging** (if not talking about AVO) consists in *sources* and receivers *collocated*  on a flat horizontal surface ("*datum*")
- **However, in reality, we** record at finite *sourcereceiver offsets*, over complex topography, and in complex nearsurface structure*.*



- In the "CMP stacking"-type imaging, **two types of corrections** are applied to compensate these factors:
	- Statics "place" sources and receivers onto the datum;
	- Normal Moveout Corrections "transforms" the records ð. into as if they were recorded at collocated sources and receivers.
- **As a result** of these corrections (plus stacking to attenuate noise), we obtain a stacked *zero-offset section*
- Another type of imaging is "pre-stack", which obtains the final image directly from the full source-receiver geometry

# Normal moveout



- Symmetric hyperbola
- Reflected rays propagate as if from a source at depth





### Normal moveout

- The NMO hyperbola can be approximated by a parabola
	- Works for small reflection angles (up to about 30°):

**Normal MovedUL**  
\n• The NMO hyperbola can be approximated by a parabola  
\n• Works for small reflection angles (up to about 30°):  
\n
$$
t(x) = \sqrt{t_0^2 + \left(\frac{x}{V}\right)^2} = t_0 \sqrt{1 + \left(\frac{x}{t_0 V}\right)^2} \approx t_0 \left[1 + \frac{1}{2} \left(\frac{x}{t_0 V}\right)^2\right] = t_0 + \frac{1}{2t_0} \frac{x^2}{V^2}
$$
\nQuadratic function -  
\nparabola in the (*x*,*t*) plane  
\n• To understand this and other formulas below, not  
\nthat for small  $\alpha \ll 1$ :  
\n
$$
(1 + \alpha)^n \approx 1 + n\alpha
$$

Quadratic function -> parabola in the (*x*,*t*) plane

■ To understand this and other formulas below, not that for small  $\alpha$  << 1:

$$
(1+\alpha)^n \approx 1+n\alpha
$$

### Reflection travel-times (*Multiple layers*)

For multiple layers,  $t(x)$  is no longer hyperbolic:



For practical applications (near-vertical incidence,

$$
pV_{i}<<1), \ t(x) \text{ still can be approximated as:}
$$
\n
$$
x_{n}(p) = \sum_{i=1}^{n} \frac{h_{i}pV_{i}}{\sqrt{1-(pV_{i})^{2}}} \approx p\sum_{i=1}^{n} h_{i}V_{i} \left[1 + \frac{1}{2}(pV_{i})^{2}\right] \approx p\sum_{i=1}^{n} h_{i}V_{i}
$$
\nhence: 
$$
p = \frac{x_{n}(p)}{\sum_{i=1}^{n} h_{i}V_{i}}
$$
\n
$$
t_{n}(p) = \sum_{i=1}^{n} \frac{h_{i}}{\sqrt{1-(pV_{i})^{2}}} \approx \sum_{i=1}^{n} \frac{h_{i}}{1 + \frac{1}{2}(pV_{i})^{2}} = t_{0} + \frac{1}{2}p^{2}\sum_{i=1}^{n} h_{i}V_{i}
$$

 $\mathbf{r}_n(p)$ 

 $p = \frac{n_n (P)}{n}$ 

hence:

hence: 
$$
p = \frac{\sum_{i=1}^{n} h_i V_i}{\sum_{i=1}^{n} h_i V_i}
$$

$$
t_n(p) = \sum_{i=1}^{n} \frac{h_i}{V_i \sqrt{1 - (pV_i)^2}} \approx \sum_{i=1}^{n} \frac{h_i}{V_i} \left[ 1 + \frac{1}{2} (pV_i)^2 \right] = t_0 + \frac{1}{2} p^2 \sum_{i=1}^{n} h_i V_i
$$

$$
t_n(x) \approx t_0 + \frac{1}{2t_0} \left(\frac{x}{V_{RMS}}\right)^2 \qquad \qquad \boxed{\qquad}
$$

here,  $V_{RMS}$  is the RMS (root-mean-square) velocity:



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## Normal Moveout (NMO) correction

NMO correction transforms a reflection record at offset  $x$  into a normal-incidence  $(x = 0)$  record:

$$
t(x) \to t_0 = \sqrt{t^2(x) - \left(\frac{x}{V}\right)^2} \approx t(x) - \frac{1}{2t(x)} \left(\frac{x}{V}\right)^2
$$
  
"Stacking velocity"



- Stacking velocity is determined from the data, as a parameter of the reflection hyperbola that is best aligned with the reflection event
- Note that NMO correction affects the shallower and slower reflections stronger
	- This is called "*NMO stretching"*

# Dipping reflector



- ◼ Hyperbola of the same shape but with the apex shifted up-dip
- Asymptotically the same moveouts





# Dip moveout

For small offsets  $(x \ll h)$  and dips  $(h \sin \xi \ll x)$ :



Reflector dip  $\xi$  can be measured from the *dip moveout*:

$$
\sin \xi \approx \frac{V}{2} \frac{t(x) - t(-x)}{x} \equiv \frac{V(t_{Downdip} - t_{Updip})}{2} \left\{ \frac{\text{This ratio is}}{\text{also called}} \right\}
$$

#### GEOL483. Dip moveout in CMP gathers

- The travel-time hyperbola becomes **symmetrical**
- Reflection points are **smeared** up-dip with increasing offset
- ◼ Asymptotic velocities *are greater* than the true velocity





## Stacking velocity in the presence of dip



■ For a fixed  $x_m$ , the dependence of the S-R time on the offset *x* is

For a fixed S (preceding case),  
\n
$$
\Rightarrow
$$
 smallest time (hyperbola apex)  
\nwould be for receiver  
\nlocated here  
\n
$$
f(x) = \frac{1}{V} \sqrt{(x + 2h \sin \xi)^2 + (2h \cos \xi)^2}
$$
\n
$$
f(x) = \frac{1}{V} \sqrt{(x + (2x_m - x) \sin^2 \xi)^2 + [(2x_m - x) \sin \xi \cos \xi]^2}
$$
\n
$$
f(x) = \frac{1}{V} \sqrt{(x + (2x_m - x) \sin^2 \xi)^2 + [(2x_m - x) \sin \xi \cos \xi]^2}
$$
\n
$$
f(x) = \frac{1}{V} \sqrt{(2x_m \sin \xi)^2 + (x \cos \xi)^2}
$$



# CMP Stacking velocity in the presence of dip (*cont.*)

This equation describes a hyperbola similar to the NMO equation (compare to:

): 2 2  $\mathcal{X}$  ) ). *NMO*  $\left(\frac{\lambda}{\sigma}\right)$  –  $\sqrt{\frac{1}{0} + \frac{1}{\sigma}}$  –  $\sqrt{\frac{1}{\sigma}}$  $x$ <sup>2</sup> *V* )  $\cdot$ 

Zero-offset time ( ) ( ) ( ) <sup>2</sup> <sup>2</sup> 1 cos 2 2 2 sin 2 sin cos *<sup>m</sup> m x x t x x x V V V* = + = + ( ) *t x t* = + Hyperbolic moveout

Thus, because of the dip, the effective velocity is increased:

$$
V_{Dip} = \frac{V}{\cos \xi}
$$

- This means that when stacking velocities are measured from a CMP gather, dipping reflectors will result in higher velocities (flatter reflection hyperbola)
- As a result, reflectors with conflicting dips cannot be NMO-corrected and stacked accurately.
	- Processing step called **DMO** (dip moveout correction) corrects this problem. We will talk about this method in the next lectures